

۱ **Unveiling the Pasqale Avalanche: A Perilous Prehistoric Landslide Proximity to Tehran's**
۲ **Metropolis in Darband Valley, Iran**

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۱۵ **Abstract**

۱۶ This research introduces the Pasqale rock/debris avalanche, an occurrence of a prehistoric
۱۷ landslide in the upstream region of Darband Valley in northern Tehran metropolitan that is
۱۸ exemplified as an instance of a high-risk geohazard encountered in the proximity of a
۱۹ metropolitan area. Its initiation as a slide from a steep scarp, with an elevation of approximately
۲۰ 3000 m, nestled within Eocene volcanic rocks and tuff, is noted. The local geology and
۲۱ geomorphology of the Pasqale landslide are described in this paper, relying on topographical
۲۲ data, satellite imagery, and field observations. Through various considerations and the utilization
۲۳ of high-resolution satellite data, the total volume of the landslide is estimated to be in the range
۲۴ of 800,000 m³. Remarkably, the occurrence of this landslide is found to be influenced by intense
۲۵ fracturing and hydrothermal alterations of the Eocene pyroclastic rocks. Moreover, the seismic
۲۶ aspect of the region is emphasized, with particular attention given to the likelihood of a large
۲۷ earthquake being the most probable triggering factor for the Pasqale avalanche, originating from
۲۸ the Mosha or North Tehran faults. The significance of the cascading hazards that may be brought
۲۹ about following a major earthquake event in the northern Tehran metropolitan area is featured in
۳۰ this paper.

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۳۲ **Keywords**

۳۳ Coseismic landslide, earthquake hazard, Central Alborz, Tehran, Pasqale

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۳۵ **1- Introduction**

۳۶ Landslide is the movement of a mass of rock, earth or debris materials down a slope of a
۳۷ mountain (Cruden, 1991). Landslides are natural phenomena that can cause disruptions and damages to
۳۸ the built environment and human societies in their path. For example, landslides can damage the
۳۹ transportation routes and in mountain valleys which is a source of considerable economic losses.
۴۰ Additionally, rapid landslides can lead to large fatalities in local environments (e.g. Turner, 2018).
۴۱ Landslides constitute a significant geohazard in emerging mountainous regions globally, spanning from
۴۲ the Alp-Himalaya belt in Eurasia to the Andes in South America, as well as surrounding the mountain
۴۳ ranges of New Zealand (e.g., Glade 2003; Crozier 2005; Dortch et al. 2009; Hasegawa et al. 2009;
۴۴ Zerathe et al. 2014; Shoaei 2014; Wood et al. 2015; Aslan et al. 2020; Grima et al. 2020; Delgado et al.
۴۵ 2020). Landslide volumes (V) exhibit remarkable diversity, ranging from a few hundred to several billion
۴۶ cubic meters, with the latter being termed giant or mega landslides ($V > 10^8$ m³). Examples of Earth's
۴۷ largest terrestrial landslides include the Seimareh landslide in the Zagros (Shoaei and Ghayoumian,
۴۸ 1998), the 1792 Unzen-Mayuyama mega-slide in Japan (Wang et al. 2019), and the TsergoRi landslide in
۴۹ Nepal (Weidinger et al. 1996), among others.

۵۰ Korup et al. (2007) conducted an extensive review of giant landslides, exploring their relationship
۵۱ with topography and erosion based on the analysis of 300 cases. Their findings reveal that nearly two-
۵۲ thirds of giant landslides occurred within the steepest 5% of land surfaces, predominantly in deeply
۵۳ incised valleys, along the peripheries of active mountain belts (i.e., fault-bounded fringes), and in
۵۴ volcanic arcs. These distributions align with regions characterized by elevated long-term erosion rates (~
۵۵ 4 mm yr⁻¹), affirming the role of giant landslides in expedited mountain denudation. Significantly, most

eroded landslide deposit volume is concentrated in the slightest yet steepest sections of mountain belts and volcanic arcs.

Pánek et al. (2016) identified giant landslides in the Caspian Sea region of western Kazakhstan, where more than 100 slope failures mobilized volumes exceeding 10^8 m^3 , even along basal failure planes with gradients as low as $\sim 5^\circ$. These events primarily occurred during Pleistocene Caspian Sea-level high stands or during the Holocene, as substantiated by C_{14} dating, and were characterized by lateral rock spreads involving competent limestones covering weak claystone beds.

Rowberry et al. (2023) recently compiled a database of giant landslides on volcanic islands, expanding our understanding of these phenomena. Pánek et al. (2012) introduced the seismic-induced giant Uspenskoye landslide in the Kuban River valley in the Northern Caucasus (Russia). This landslide is believed to have originated catastrophically as a rotational block slide combined with earthflow. Subsequent Kuban River erosion of the landslide toe triggered secondary collapses, resulting in multiple rotational landslides. Strom and Wang (2022) undertook a comprehensive review of rockslides in Central Asia, describing several giant rock slides in the Pamir, Tien Shan, and Djungaria mountains triggered by large historical and instrumental earthquakes.

Despite the long history of geological investigations, giant landslides only sometimes present clear and readily identifiable features to experts. The Alborz Mountains in northern Iran constitute a $\sim 2000 \text{ km}$ long belt of folded Paleozoic to Cenozoic rocks, connecting the Caucasus and Hindukush mountains to the northwest and northeast, respectively. This polyorogenic folded belt, evolving during Cimmerian and Alpine orogeny, serves as the focus of active deformation and has witnessed significant historical and instrumental earthquakes.

Historical accounts and recent studies underscore the pivotal role of coseismic land sliding in the Alborz Mountains, affecting both the belt's geomorphological evolution and its socioeconomic landscape. Nevertheless, many landslides in the Alborz Mountains still need to be studied, warranting further efforts to establish a comprehensive database. This paper introduces the Pasqale giant landslide, situated in the upper slopes of the Darband Valley in the north of the Tehran metropolitan area, for the first time

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83 2- Methods

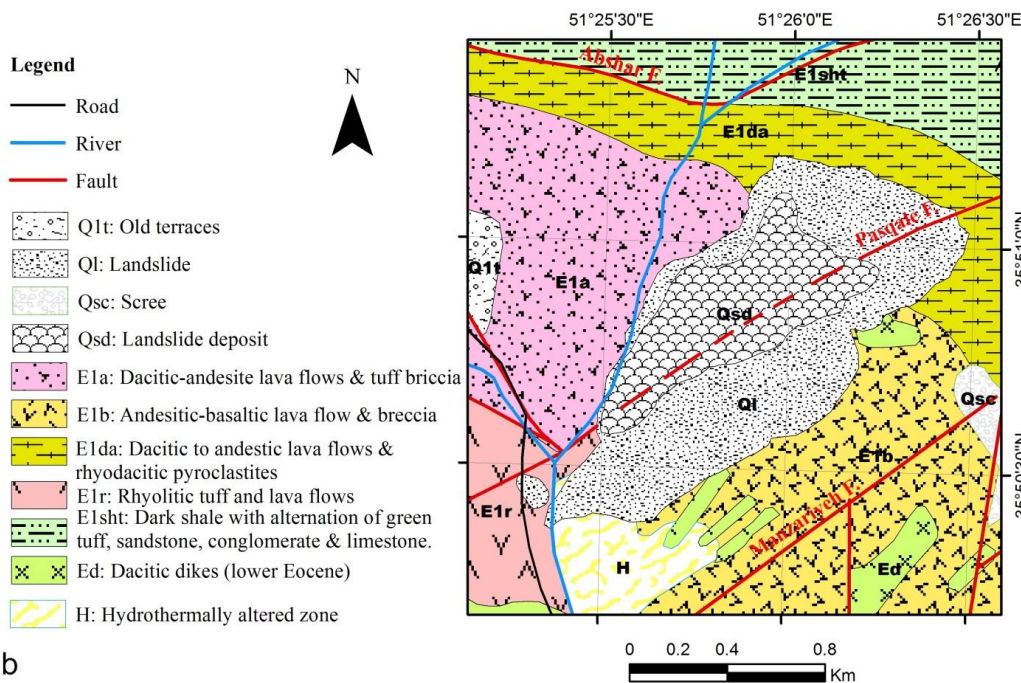
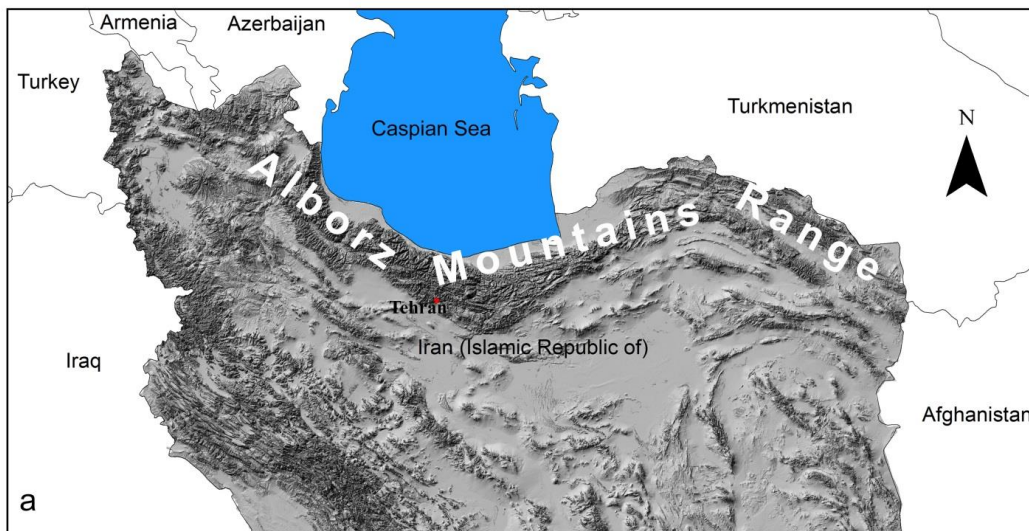
84 Given the Pasqale landslide geohazard significance and proximity to Tehran, this study offers preliminary 85 insights derived from field observations, satellite imagery, and GIS analysis. Measurement on the area 86 was based on field observations and analysis of high resolution satellite data from Google Earth, Bing 87 data and Sentinel satellite data. Topographic profiles and measurements was done by preparing a Digital 88 Elevation Model (DEM) of the slide zone that was generated using 1:25,000 topographic data with a 89 contour interval of 10 m (Figures 1&2)

90 3- Observations

91 3-1- Geological setting and geomorphology of Pasqale landslide

92 The Pasqale landslide is situated within the Darband Valley's upper reaches, located north of 93 Tehran, in the southern flanks of the Central Alborz Range (Figure 1a). The geological formations 94 exposed in the entire landslide region and its associated deposits are belonging to the Eocene Karaj 95 Formation. Figure 1b illustrates the broader geological context of the study area, adapted from the 96 1:100,000 geological map of Tehran (Emami et al. 1993). The upper levels of the slide zone, comprising 97 the scarp and an exposed rupture surface, consist of dacitic to andesitic lava and rhyodacitic pyroclastic 98 rock types (E1da). In the middle section of the landslide zone, the composition shifts to include dacitic to

۹۹ andesitic lava with tuff breccia (E1a) and andesitic to basaltic lava flow and breccia (E1b). Notably, these
 ۱۰۰ rock units exhibit varying degrees of alteration to the south of the slide zone.



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 ۱۰۲ Figure 1. a) SRTM shaded relief map showing the general trend of the Alborz Mountains Range in the
 ۱۰۳ north of Iran. The red dot shows the location of the area of study; b) Geology map of Pasqale study area
 ۱۰۴ in the northern Tehran. The landslide district in the map includes volcanic rocks, green tuffs and shales of
 ۱۰۵ Eocene Karaj Formation. The geology map is modified after Emami et al. (1993).

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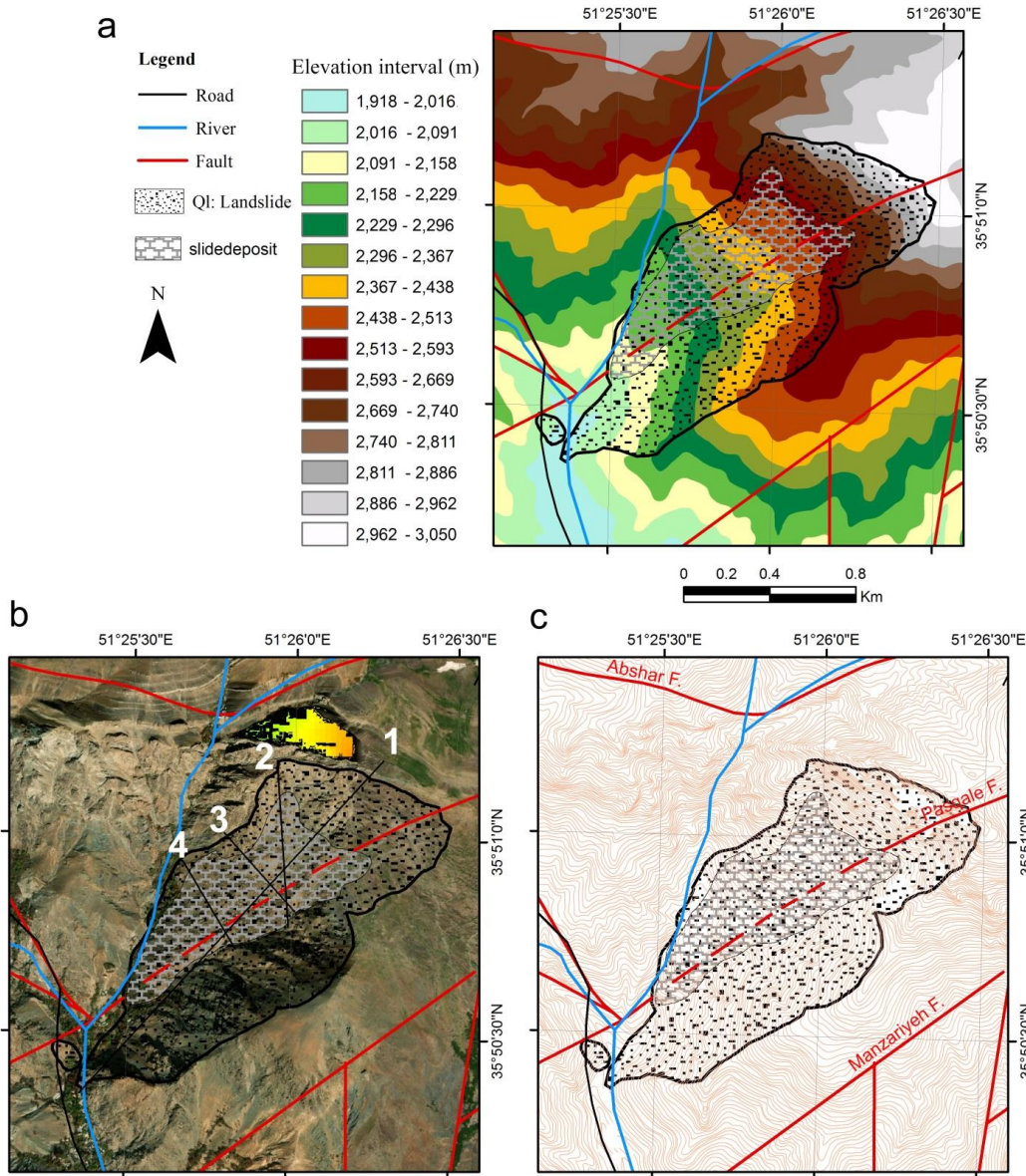
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The topographic contrast between the upper levels of the landslide scarp (>3000 m) and the current level of the slide mass is about 1000 m. Figure 2b illustrates the spatial position of the slide zone with major fault strands and the location of four lines along with topographic profiles. Figure 2c provides the original topographic map employed for making DEM. The plotted topographic profiles encompass one longitudinal and three transverse profiles traversing the slide zone (Figure 3). The longitudinal profile (profile 1 in Figure 3) reveals a nearly vertical slope on the higher levels of the slide zone, corresponding to the prominent Pasqale landslide scarp, with the surface of the rupture only partially exposed. The transverse profiles 3 and 4 (Figure 3) capture secondary topographic undulations attributable to the slide body, inclusive of slide deposits.

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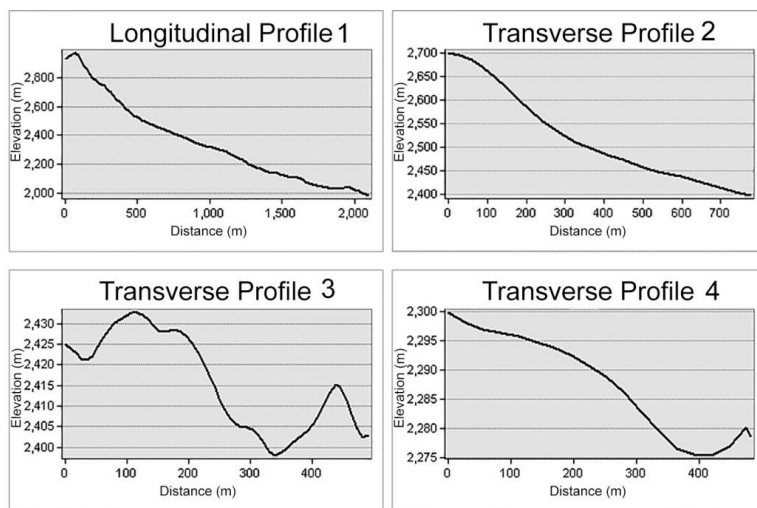
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Figure 2. a- Digital elevation model (DEM) of the Pasqale landslide and surrounding area study area that produced from 10 m intervals topographic map using GIS. Boundary of landslide zone and its deposit with major faults have marked on the DEM; b- satellite image of the same area (Bing Images, 2023) on which the location of four topographic profiles (in Figure 3) is shown by white lines. c- Original topographic data that was used for DEM generation (achieved from National Cartographic Center of Iran).

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Figure 3. Topographic profiles across the Pasqale landslide based on DEM (10 m); profile 1 is longitudinal and 3 others are transverse. Locations of topographic profiles are shown in Figure 2b by white lines. The horizontal axes of profiles shows the distance (m) from the starting points of profiles that marked by their profile number in Figure 2b. In Profile 1 note to the higher slope gradient in the upper scarp; the topographic difference between transverse profiles (profiles 2, 3 and 4) is mainly due to accumulation of the slide deposit.

Comment [a3]: Need revision

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While Figures 1b and 2 delineate major fault strands in the study area, Figures 4a and 4b depict a densely fractured system within the exposed rocks of the uppermost segment of the original rock mass. Figure 4a provides an on-site view of the primary scarp (MS) and the exposed surface of rupture (SOR) of the Pasqale landslide. Conversely, Figure 4b presents an annotated satellite image of the region (Bing, 2023), revealing densely fractured rock units in the northern portion. The Pasqale landslide was initiated from the north and northeastern slopes of the Darband Valley, as indicated by the arrows in Figure 4b. However, it remains reasonable that the Darband Valley may have experienced prior landslides from other slopes, necessitating a geophysical survey to ascertain basal rock mass variations beneath the current Pasqale slide deposits.

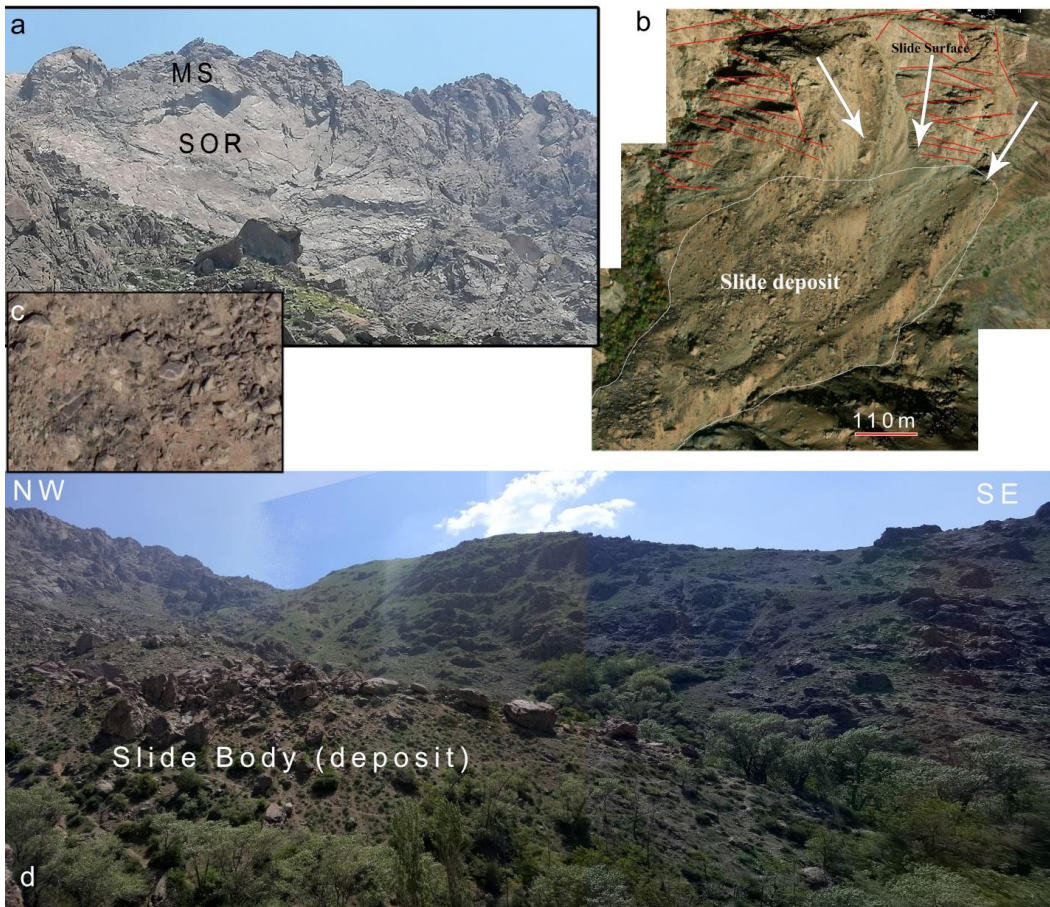
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3-2- Rough estimation of landslide volume

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Estimating the volume of the Pasqale landslide is challenging due to the absence of original topographic data predating the landslide event and limited subsurface information from the slide deposit.



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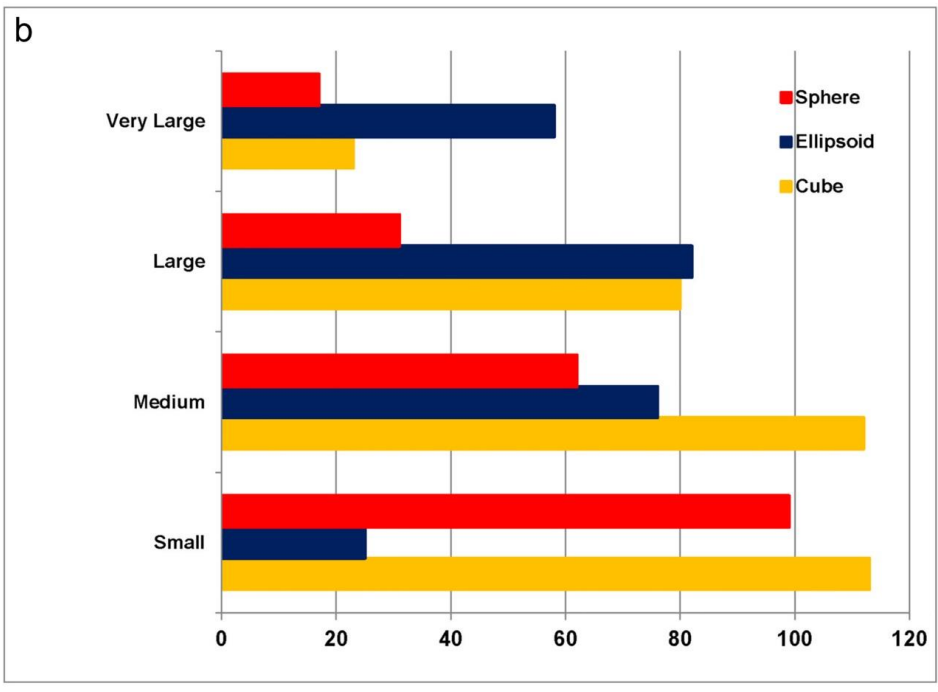
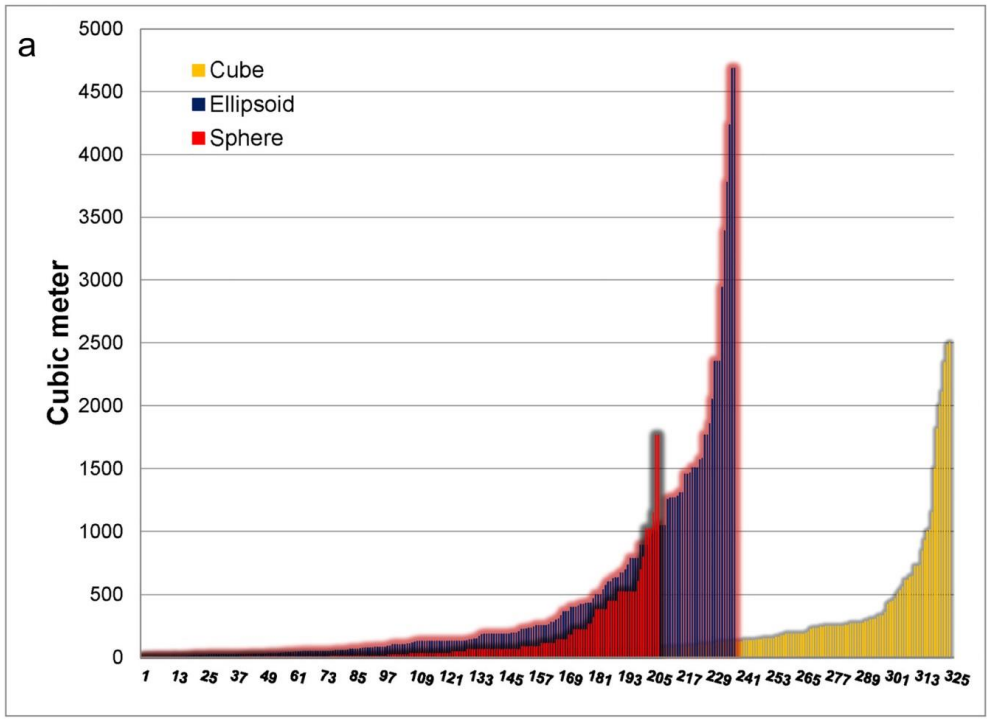
Figure 4.a- Field photograph of the Main scarp (MS) and Surface of Rupture (SOR) of the Pasqale landslide; topographic height at the top of main scarp is ~2980 m.; b- photomap of the Pasqale landslide prepared by the satellite image from Google Earth (2023) that shows the boundary of slide deposit (white polygon), movement direction (white arrows) and slide surface c- a sample of satellite image with high spatial resolution provided by Bing (2023) that was used for measuring the size of rock blocks; d- field photograph showing the panorama view of slide body in which rock blocks are variable is size. The slide deposit is bounded by tall white poplar trees Sepidar in Persian, *Populus alba*.

Comment [a4]: Need revision

Nonetheless, to gain a preliminary rough estimation of the landslide volume, a volumetric measurement of exposed rock blocks on the surficial slide deposit was conducted using high-resolution satellite images and GIS (Figures 4c and 4d). By this approach, rock blocks are classified as cube, spherical, or ellipsoid shapes. While satellite images are inherently two-dimensional, certain assumptions were made to estimate rock block volumes. Specifically, measurements were taken for the long and minor axes for cube and ellipsoid shapes, with the third axis considered equivalent to the middle axis. For spherical rock blocks, only the radius was necessary for volume calculations. While this method is approximated, it provides an estimate of rock block volumes. Figure 5a illustrates the share and volume size distribution of measured rock blocks within the Pasqale landslide deposit categorized by their shape

165 (spherical, cube, and ellipsoid). Figure 5b offers comparative histograms of volume size distributions for
166 the different rock block types. The data reveals that cube-shaped rock blocks predominate in terms of
167 occurrence, whereas ellipsoid shapes contribute to larger blocks. Classification by size indicates that
168 small to medium cubes and small spherical rock blocks are more common. In contrast, ellipsoid rock
169 blocks feature more prominently in the larger block category.

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1971 Figure 5. a- shape and volume size distribution of measured rock blocks of Pasqale landslide deposit by
1972 their type (spherical, cube and ellipsoid); b- comparative histogram of volume size distribution of
1973 measured rock blocks.

Comment [a5]: Shape and volume size

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1976 The roughly estimated volume of surficially exposed rock blocks within the Pasqale landslide is
1977 approximately $\sim 200,000 \text{ m}^3$. Currently, the timing of the Pasqale landslide remains undetermined, and the
1978 extent of fine-grain slide deposits that fluvial processes may have eroded remains uncertain. However,
1979 various observations, such as the presence of thriving vegetation around the deposit (Figure 4d), the
1980 distinct surface rupture, and known cascades along the flowing Darband River, suggest that the landslide
1981 likely occurred during the Holocene. Nevertheless, it is reasonable to assume that a substantial portion of
1982 the slide deposit has been eroded, while a significant proportion of fine-grained deposits remain. The
1983 erosion time series of landslide deposits is related to several factors such as climate, local topographic
1984 condition, lithology, size and cohesion of slide mass, proximity to river and fluvial system etc. the true
1985 measurement of this time series requires profound interdisciplinary studies on landscape evolution with
1986 geochemical analysis, dating and provenance or fingerprinting studies of slide deposits (e.g. Chang and
1987 Zhang, 2010; Gan et al., 2018; Del Vecchio et al., 2018; Chen et al., 2022). However, in various cases for
1988 landslides deposits are dated back to several hundreds to more than 1000 years ago, between 10 to 30% of
1989 slide deposits still remained uneroded (Del Vecchio et al., 2018; Chen et al., 2022; Koshimizu and Uchida,
1990 2023). Consequently, the estimated volume of the debris mass may be between four to six times the
1991 measured volumes of rock blocks. Accordingly, the total roughly estimated volume of the Pasqale
1992 landslide ranges from $800,000$ to $1,400,000 \pm 200,000 \text{ m}^3$. Of course, this amount is the total volume
1993 before the start of erosion. Therefore, it is possible that the current value of the volume of the landslide
1994 mass is the same number as $800,000$ cubic meters.

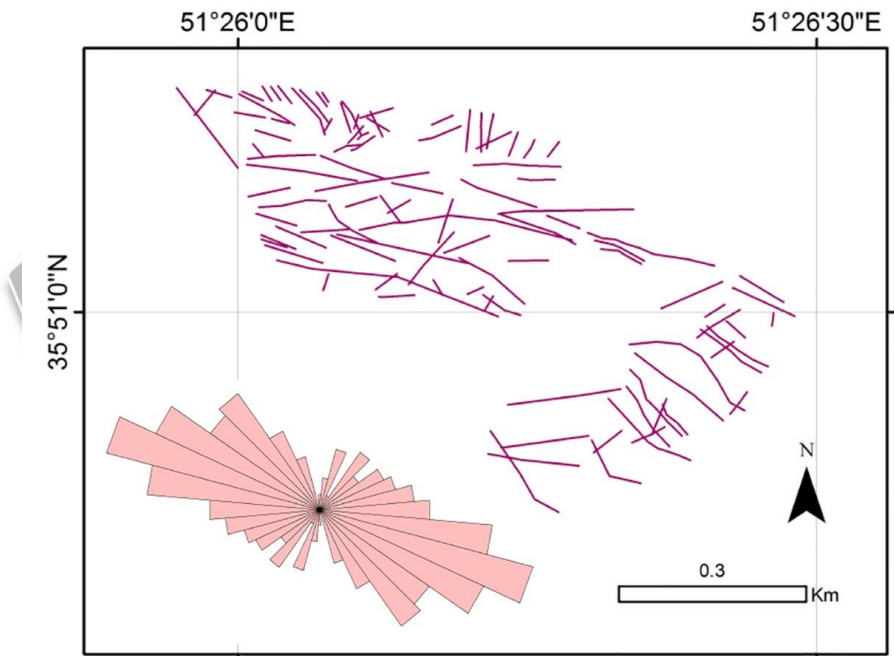
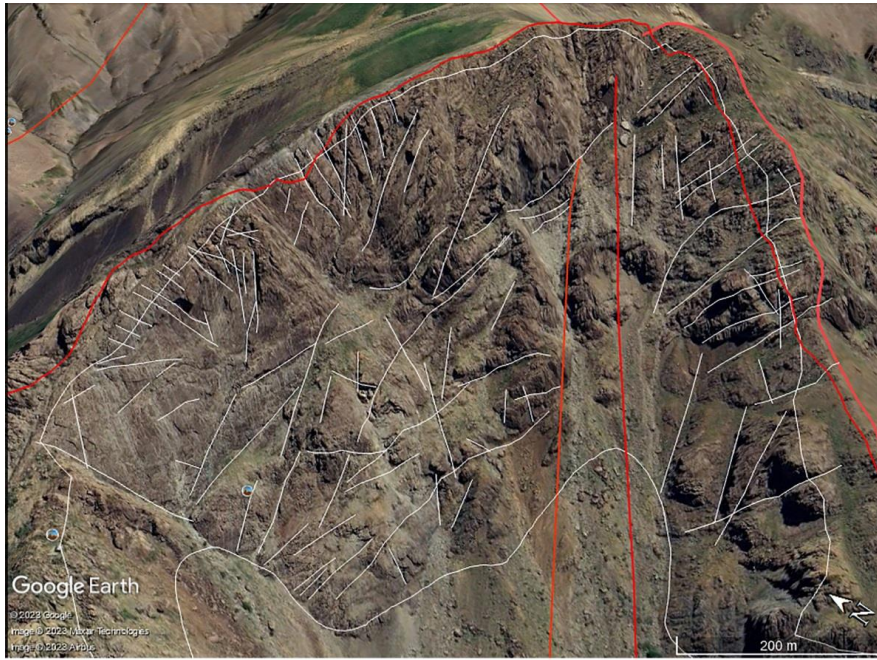
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Of course, this amount is the total volume before the start of erosion. Therefore, it is possible that the current value of the volume of the landslide mass is the same number as $800,000$ cubic meters.

1995 4- Discussion

1996 The Pasqale landslide, introduced for the first time in this study, is a substantial landslide situated
1997 within the Darband Valley on the southern slopes of the Central Alborz Mountains. The slide zone
1998 primarily comprises weathered and altered volcanic rocks and tuff belonging to the Eocene Karaj
1999 Formation. Notably, the slope gradient between the slide scarp and the toe is remarkably steep, with an
2000 elevation difference of approximately 1000 m . The Darband Valley may have a history of older
2001 landslides, thus representing a paleolandslide zone that warrants further investigation. Extensive
2002 fracturing of higher slopes of landslide area provided required condition for disintegration of rock blocks,
2003 where the several sets of fractures intersect with each other (Fig. 6).



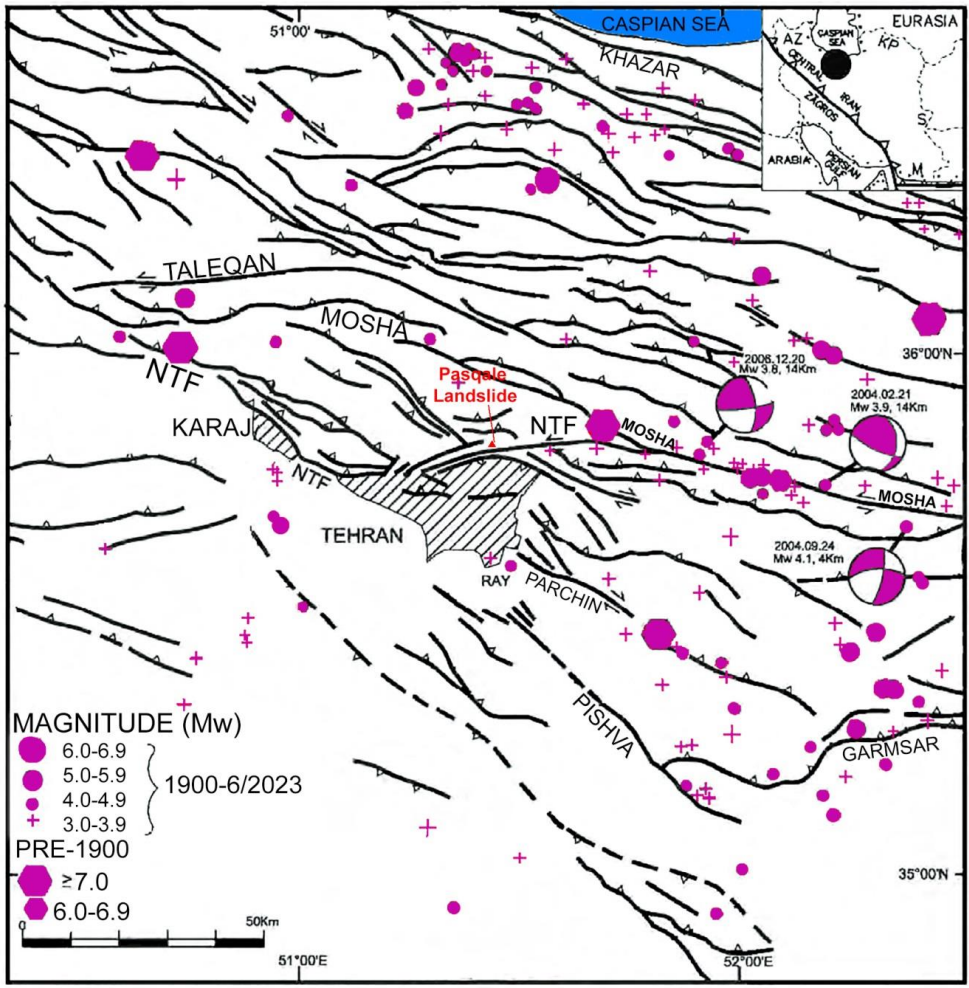
۲۰۴
 ۲۰۵ Figure 6. More detailed map of fracture distribution on the top portion of Pasqale landslide with their rose
 ۲۰۶ diagram showing several sets of available fractures that are more frequently ~E-W. Intersection fracture
 ۲۰۷ sets provided suitable condition of rock blocks to destabilize.

۲۰۸ At this stage, we estimate the volume of the Pasqale landslide to range from 800,000 to 1,400,000
۲۰۹ $\pm 200,000 \text{ m}^3$. While the Pasqale landslide does not meet the criteria of a giant landslide, it is still
۲۱۰ reasonably classified as a large-scale landslide. Based on observations of the exposed sliding surface, the
۲۱۱ high steep gradient, and the overall morphology, it is inferred that the Pasqale landslide was likely
۲۱۲ initiated as a rock-debris slide. However, it experienced a transition to a rock/debris avalanche due to the
۲۱۳ following key observations:

- ۲۱۴ 1. The exposed slide surface discontinues downward (Figure 4a), indicating that the slide body is
۲۱۵ entirely detached from the slide surface after minimal movement.
- ۲۱۶ 2. Of greater significance, rock blocks of considerable size, comparable to the dimensions of a room
۲۱۷ or small building and evaluating several hundred tons, were pushed several kilometers from their
۲۱۸ original positions. Such occurrences align with the characteristics of rock avalanches, which are
۲۱۹ recognized globally as coseismic landslides (e.g., Keefer 1984; Jibson et al. 2006; Dunning et al.
۲۲۰ 2007). Rock avalanches represent one of the most common fatal landslide types in the Alborz and
۲۲۱ Zagros regions (Roberts and Evans 2013; Ehteshami-Moinabadi 2019; Gutiérrez et al. 2023).

۲۲۲ As previously mentioned, Korup et al. (2007) identified that nearly two-thirds of giant landslides
۲۲۳ are concentrated within the steepest 5% of land surfaces, particularly in deeply incised valleys, along the
۲۲۴ peripheries of active mountain belts (fault-bounded fringes), and in volcanic arcs. The conditions
۲۲۵ prevailing in the Darband Valley, characterized by its steep and deeply incised terrain, align with the
۲۲۶ criteria observed for giant landslides.

۲۲۷ From an active tectonic and seismotectonic perspective, the Mosha and North Tehran faults
۲۲۸ represent two major seismic faults that have experienced historical and pre-historical large earthquakes
۲۲۹ ($M > 7.0$) (Fig 7). These faults are located north and south of the study area, respectively (e.g., Berberian
۲۳۰ and Yeats 1999, 2001, 2017; Ritz et al. 2012; Talebian et al. 2016). While a detailed discussion of the
۲۳۱ seismotectonic aspects of the region is beyond the scope of this paper due to length constraints, interested
۲۳۲ readers are directed to the aforementioned references and additional works such as Berberian et al.
۲۳۳ (1985), Landgraf et al. (2009), Solaymani-Azad et al. (2011), Ghassemi et al. (2014), Solaymani-Azad
۲۳۴ (2023), Ehteshami-Moinabadi and Nasiri (2023) and Ehteshami-Moinabadi et al., (2023).



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 ۲۳۶ Figure 7. Map showing the historical and instrumental seismicity of the Tehran region and surrounding
 ۲۳۷ area; instrumental seismic data achieved from Iranian Seismological Center. Historical seismicity and
 ۲۳۸ faults based on Berberian and Yeats (2017). The location of Pasqale Landslide is marked by red triangle.

۲۳۹ Given the well-established relationship between earthquakes and large landslides in the Alborz
 ۲۴۰ Mountains (e.g., Berberian et al. 1992; Asadi and Zare 2014; Ehteshami-Moinabadi and Nasiri 2019), and
 ۲۴۱ considering the proximity of the North Alborz and Mosha faults to the Pasqale landslide, it is reasonable
 ۲۴۲ to assert that an earthquake represents the most probable triggering mechanism for this landslide. Further
 ۲۴۳ research and investigation are essential to refine our understanding of the specific seismic conditions and
 ۲۴۴ factors that led to the initiation of the Pasqale landslide.

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 ۲۴۶ **5- Conclusion**

۲۴۷ In conclusion, the Pasqale landslide represents a significant prehistoric event characterized as a
۲۴۸ large rock/debris avalanche originating from a slide in the Darband Valley's upper reaches, situated north
۲۴۹ of Tehran. The rough estimated volume of this landslide falls within the range of 800,000 to 1,400,000 ±
۲۵۰ 200,000 m³. Notably, the composition of thrown rock blocks varies in size, ranging from a few cubic
۲۵۱ meters to several hundreds of cubic meters, with some exceptionally large blocks.

۲۵۲ The geological components contributing to this event primarily comprise volcanic rocks and tuff,
۲۵۳ characterized by intense fracturing. Additionally, the upper surface of the landslide exhibits a near-
۲۵۴ vertical slope. Moreover, the Darband Valley exhibits features consistent with the potential occurrence of
۲۵۵ paleolandslides, warranting further investigation.

۲۵۶ From a seismotectonic viewpoint, the study area is identified as an active area. In this context, the
۲۵۷ most plausible triggering factor for the Pasqale landslide is a significant earthquake originating from
۲۵۸ either the Mosha or North Tehran faults. This observation underscores the significance of the Pasqale
۲۵۹ landslide as an illustrative example of the hazards associated with coseismic mass movements in the
۲۶۰ northern vicinity of the Tehran metropolitan area. The implications of such events on the safety and
۲۶۱ preparedness of this densely populated urban region are of paramount concern, highlighting the need for
۲۶۲ continued research and assessment in geohazard mitigation and disaster management.

۲۶۳ **6- Declarations**

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۲۶۷ **Author contributions**

۲۶۸ All authors contributed to the study conception and design. Material preparation, data collection and
۲۶۹ analysis were performed by [Mohsen Ehteshami-Moinabadi] and [Elahe Feizabadi]. The first draft of the
۲۷۰ manuscript was written by [Mohsen Ehteshami-Moinabadi] and all authors commented on previous
۲۷۱ versions of the manuscript. Shahram Nasiri read the draft and made contribution on final revisions. The
۲۷۲ original idea about this landslide was inspired by Mohsen Ehteshami-Moinabadi and Shahram Nasiri. All
۲۷۳ authors read and approved the final manuscript.

۲۷۴ **Conflicts of interests (CEI)**

۲۷۵ The authors declare that they have no known competing financial interests or personal relationships that
۲۷۶ could have appeared to influence the work reported in this paper. The authors have no conflicts of interest
۲۷۷ to declare. All co-authors have seen and agree with the contents of the manuscript and there is no
۲۷۸ financial interest to report. We certify that the submission is original work and is not under review at any
۲۷۹ other publication.

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۲۸۵ **8- References**

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